

tively. Arrows show approximate width of the belt, and partly overlapped strips, photographed during one revolution. Number pairs designate all frames included

Ridge marking rim crest of highly

Area in which dark craters are concentrated

degraded or buried crater

GEOLOGIC MAP OF THE MACROBIUS QUADRANGLE OF THE MOON
By
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The Macrobius quadrangle is in the northeast quadrant of the Moon's near side. Although predominantly a highland area centered around the Taurus Mountains (Montes Taurus), it is bounded by three major mare-filled basins: Tranquillitatis, the oldest, to the south; Serenitatis to the west; and Crisium on the east. Most of the geologic units within the quadrangle have been profoundly affected by the formation of these basins and by that of the younger Imbrium basin farther to the northwest. Some of the youngest material on the Moon, believed to be of volcanic origin, blankets part of the mare and terra along the southwest margin of the map. Sampling this material will be a primary objective of the Apollo 17 mission, whose proposed landing site is approximately at lat. 20° 10′ N., long. 30° 45′ E. near the western edge of the map area.

STRATIGRAPHY

GEOLOGIC SUMMARY

GEOLOGIC ATLAS OF THE MOON

MACROBIUS QUADRANGLE

I-799 (LAC-43)

Crater materials are dated relatively and assigned positions in the lunar stratigraphic column primarily on the basis of their morphology, using guidelines and criteria set forth by Pohn and Offield (1970) and Trask (1971). Stratigraphic units which mark the base of the Imbrian System cannot be confidently recognized within the quadrangle. However, superposition and embayment relationships between classified crater materials and regional units such as mare, plains, and some terra deposits are used to determine the position of the mapped units relative to the Imbrian—pre-Imbrian boundary. Pre-Imbrian.-Pre-Imbrian materials include those of craters (pIc), massifs (pItm), and parts of hilly (IpIh) and rugged (IpIr) terra units not separable on the basis of morphology from their Imbrian components. Most pre-Imbrian craters can be recognized by their outlines, but their materials are indistinguishable from the surrounding terrain, especially where they occur within the hilly unit. Some of these craters predate the adjacent basins, except possibly Tranquillitatis; their advanced stage of degradation may be due to erosion and partial burial by basin ejecta. Many craters, however, postdate the basins, and the extreme modification of their rims and walls is believed to be primarily the result of cumulative structural deformations

accompanied by mass wasting. The massif material resembles that of the Apennine Mountains near the Apollo 15 landing site (Carr and El Baz, 1971) and probably also consists of breccias ejected from adjacent basins. Episodic faulting and uplift subsequent to basin formation have exposed the breccia deposits as high smooth scarps, kept fresh and bright by continuous downslope movement of material. The massif unit locally is overlain by, or grades upward into, hilly material. The hilly material and the rugged terra unit form a complex believed to be made up of basin and crater ejecta. Morphological differences between the two units may reflect distance from basin of origin and relative age. The more distal ejecta facies (IpIr) is smoother and resembles the Fra Mauro Formation, whereas the hilly unit (IpIh) is closer to its basin(s) of origin and more nearly resembles the Alpes Formation (Page, 1970). Thus, the three units represented by massif, hilly, and rugged materials, though morphologically distinguishable, may consist of lithologically similar breccias whose emplacement ages are mostly pre-Imbrian, like their basins of origin; admixtures of Imbrium-basin, Imbrian-crater, and mass-wasted materials account for the Imbrian to pre-Imbrian age range assigned to the hilly and rough terrain units. Imbrian System.—Terra, plains, and mare materials generally occupy successively lower topographic levels in the quadrangle. Terra material (It) is a blanket which partly fills depressions and subdues subjacent terrain without exhibiting any particular morphologic characteristics of its own. It is chiefly a mixed unit made up of reworked surficial materials, impact ejecta, mass-wasted products, and possibly some volcanic materials locally introduced. As the terra unit is formed by processes which have continued in varying degrees through-

out lunar history, its overall age depends largely on when these processes or events culminated in any particular region. In the Macrobius quadrangle the unit appears to have been mostly developed during Imbrian time. Mare units (Im<sub>1</sub>, Im<sub>2</sub>) are typical of those elsewhere on the Moon now known from Apollo samples to consist mostly of basaltic lavas. Plains material (Ip) appears to be intergradational with both the terra and mare deposits. In the Macrobius quadrangle, and as interpreted elsewhere by many previous workers, it is probably of mixed origin (for example, McCauley, 1972; Wilhelms, 1972). In some areas the unit may consist of pyroclastics or of relatively textureless impact ejecta; in other places it may consist of older and more thoroughly impact-churned lava flows extruded prior to the onset of the major pulse of mare volcanism. Descriptions of the Cayley Formation (a plains unit) made by Apollo 16 astronauts and documented by the Apollo Lunar Geology Investigation Team (1972) indicate that much of the material is composed of fragmental rocks, probably impact breccias. Hilly and furrowed terra (Ihf) occurs in many places on the Moon, and

Hilly and furrowed terra (Ihf) occurs in many places on the Moon, and sampling it was a major objective of the Apollo 16 mission at the Descartes landing site (Milton and Hodges, 1972). This type of material generally has been considered to be of volcanic origin, but it closely resembles the hilly terra unit (IpIh) and the possibility that it also is related to a basin-forming event cannot be excluded (Trask and McCauley, 1972). Preliminary descriptions of some Apollo 16 samples indicate that both crystalline rocks and breccias have been found in this material.

Secondary craters formed by ejecta from Posidonius (northwest corner of the map) are younger than the numerous rilles upon which they are superposed. The rilles, in turn, appear to transect patches of Imbrian mare materials. Posidonius, however, is embayed by both recognized mare units. These seemingly anomalous age relationships indicate that pre-existing rilles may preserve a fresh, relatively sharp appearance even where partly filled by younger lava flows, or that some mare material is older than Posidonius.

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Eratosthenian and Copernican Systems.—The darkest mare material and sharp, bright craters without rays are assigned to the Eratosthenian System. All rayed craters, including those with dark halos, are Copernican in age. Overlapping ray patterns of various intensity and large contrasts in morphology between small Copernican craters indicate a wide range in crater ages within the Copernican System. These age subdivisions are not made at the 1:1,000,000 map scale but have been attempted at the 1:250,000 (Scott and Carr, 1972) and 1:50.000 (Lucchitta, 1972) scales.

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Among the youngest Copernican units, two are of special interest and will be primary objectives at the Apollo 17 landing site. Dark mantling material of probable volcanic origin (Cd) covers mare and terra materials in the southwest part of the quadrangle. This units resembles the Sulpicius Gallus Formation mapped along the southwestern margin of Mare Serenitatis (Carr, 1966). The dark material within the Macrobius quadrangle is considered to be Copernican because (1) it has a much lower density of small bright craters than adjacent mare surfaces, and (2) the oldest superposed craters appear to be about middle Copernican in age on the basis of their morphology (Trask, 1971). Bright mantling material (Cb) in the same locality has some of the characteristics of a terrestrial landslide or debris flow and extends across a plains surface at the base of a scarp that borders a massif. It appears relatively thin, and its contact with the dark mantling material is serrate and indistinct.

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STRUCTURE

Tectonic activity throughout this region has produced prominent structural trends expressed by normal faults, ridges, depressions, crater chains, and lineaments of subtle, unresolved topographic features.

Azimuth-frequency distributions show high concentrations of the structural elements in the N. 45°-60° W. and N. 40°-50° E. directions, similar to those of

ments of subtle, unresolved topographic features. elements in the N. 45°-60° W. and N. 40°-50° E. directions, similar to those of the "lunar grid" in the northern hemisphere (Strom, 1964). Northwest azimuths are especially common, however, and closely coincide with radials to the Imbrium, Serenitatis, and Crisium basins, whose centers (and the center of the quadrangle) lie approximately along a great circle. Thus, global stresses responsible for the northwest grid sets seem to have been reinforced and reactivated by the formation of the basins. Rimae Littrow I, II and Rima Chacornac I are large grabens concentric with the main rim of Mare Serenitatis; they extend nearly continuously for more than 300 km across mare, plains, and terra from the southwest to the northwest corners of the quadrangle. In like manner, the long rilles Rimae Römer I, II and Rima Bond I are roughly concentric with the Crisium basin, as is the large mare- and plains-filled depression extending from near Macrobius A to the northeast corner of the map. Between this depression and an ill-defined arcuate trough adjacent to the Crisium rim, a broad high area, extending through the crater Macrobius, marks an outer ring of the Crisium

GEOLOGIC HISTORY

The large impacts which successively formed the Tranquillitatis, Serenitatis, and Crisium basins covered the Macrobius quadrangle with a series of overlapping ejecta blankets as much as 2.5 km thick (Short and Forman, 1970, fig. 4). This complex of interlayered breccias also received smaller contributions of impact ejecta from large craters within the area, as well as from nearby basins such as Fecunditatis and Nectaris. The ejecta blankets and underlying lunar crust were fractured and faulted along global and basin radial systems by internal stresses, including seismicity resulting from the high flux of large impacting bodies during this early time. Mass wasting and slumping, which appear to have been especially active from pre-Imbrian time through the Imbrian Period, have modified the highly structured lunar surface into rounded hills (unit IpIh) and rugged terrain (unit IpIr). The impact which formed the Imbrium basin and marked the beginning of the Imbrian Period may have covered parts of the quadrangle with a relatively thin mantle of ejecta. Some of this material may comprise parts of the smooth terra (It) and plains (Ip) units as well as the northwesterly lineated terrain north of Römer partly mapped as Fra Mauro by Wilhelms and McCauley (1971). The Imbrium impact induced faulting along pre-existing zones of structural dislocation and weakness which is manifest in this region by the system of distinct northwest-trending normal faults. During the Imbrian Period, lowviscosity lava flows began to be extruded, filling depressions and some crater floors with plains material and mare basalt. Mare volcanism and impact cratering continued at a decreasing rate during the Eratosthenian Period. Vestiges of waning volcanic activity in Copernican

time occur in the southwestern part of the map area, where a very dark mantle (unit Cd) resembling an ash-fall tuff partly covers mare and terra alike.

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